

Overview of useful spectral regions for Venus: An update to encourage observations complementary to the Akatsuki mission

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ABSTRACT

New tables to observe the planet Venus are presented with detailed information about the main spectral regions from 100 nm to 1 mm. The information hereby is updated thanks to the Venus Express legacy and recent ground-based observations, and we hope it can constitute a helpful tool for professional and amateur observers willing to support and coordinate with the JAXA's Venus Climate Orbiter/Akatsuki mission.

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1. Introduction

The planet Venus is permanently shrouded by a thick layer of clouds that appears featureless in the visible range. During about two centuries, attempts were made at describing the surface of Venus or measuring motions presumably apparent on its disk (Bianchini et al., 1996), but it was not until the beginning of the XXth century when a first useful narrow spectral band was found in ultraviolet, revealing global-scale dark features moving with a period of 4 days (Ross, 1928; Boyer & Camichel, 1961). During the following decades, new ranges for broadband photometry were discovered and allowed ground-based images and spectra to identify compounds of the atmosphere and clouds, characterize and constrain the atmospheric chemistry, observe the cloud morphology and winds at multiple vertical levels, estimate the temperatures at the surface and several altitudes of the atmosphere, or mapping the surface and measure the slow rotation of Venus' solid-globe (Taylor, 2014). During 2006–2014, ESA's orbiter Venus Express performed a complete characterization of the Venus atmo-

sphere, and some new narrow spectral bands were discovered with VIRTIS (Wilson et al., 2009).

High-quality data able to overcome the worse spatial resolution in ground-based observations is possible nowadays provided the outstanding improvement in instrumentation for telescopes, along with the revolution of low-cost cameras and "lucky imaging" within the amateur community (Mousis et al. 2014). Complementary ground based observations such as these are expected to cover time and spatial gaps in observations by JAXA's Akatsuki orbiter (Nakamura et al., 2007), which has been performing regular acquisition of images to study the atmospheric dynamics of Venus since its successful orbit insertion in December 2015. Ground based observations can also extend the spectral coverage of Venus' observations during the Akatsuki mission.

2. The spectral bands to observe Venus

Tables 1 and 2 are a summary of most of the spectral bands relevant to study the Venus day side and/or night one, covering from far-ultraviolet (FUV) to microwave wavelengths. Multiple vertical levels of the Venus atmosphere can be sensed through three basic mechanisms: reflection/absorption of the sunlight on the day side, fluorescence processes and thermal emission partially absorbed by atmospheric compounds and clouds above the emitting

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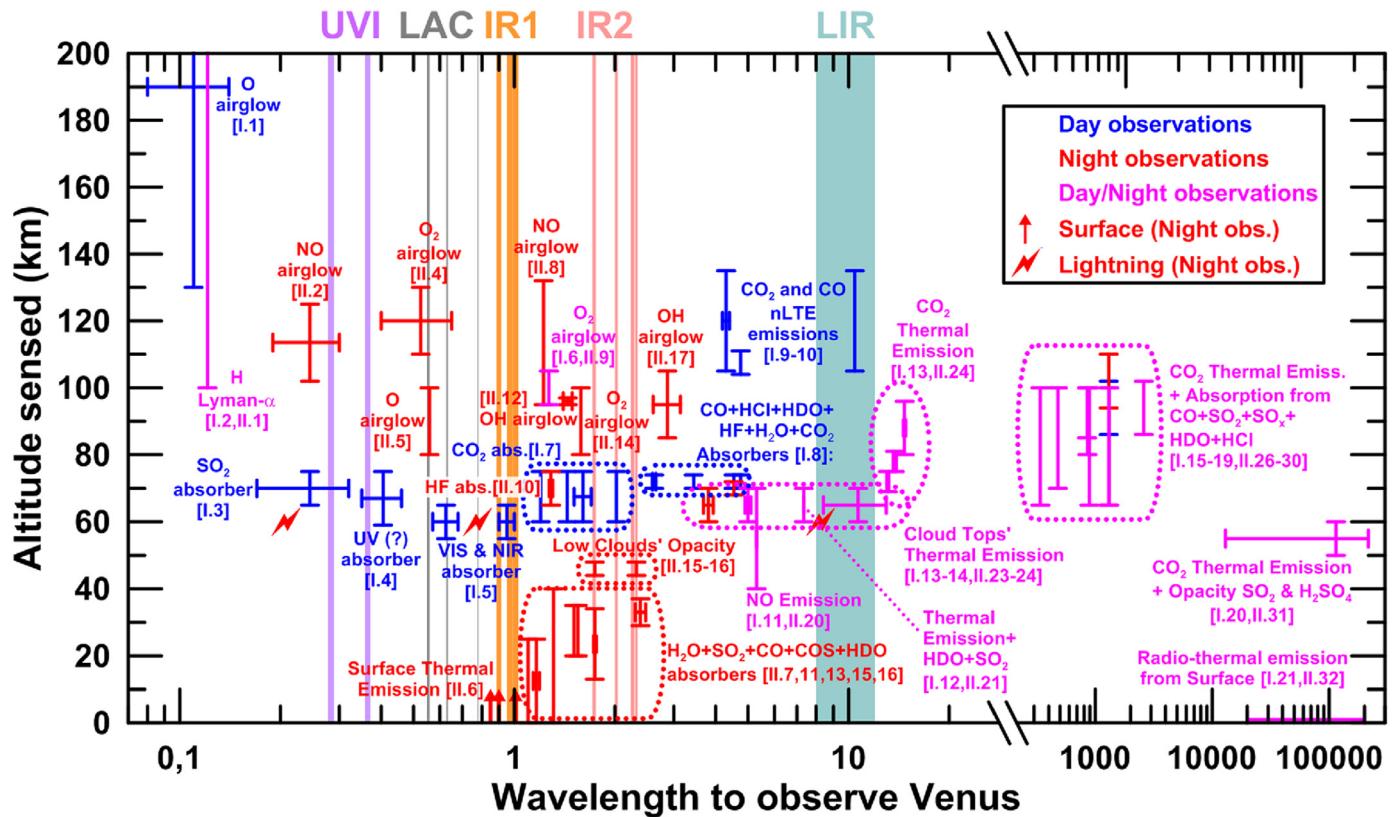


Fig. 1. Venus' spectral bands in the framework of Akatsuki's payload. References are specified in square brackets (e.g. I.19.II.30 stands for row 19 in Table 1 and row 30 in Table 2, respectively.).

layer. These processes can act simultaneously, complicating the observations; for example, the $1.27\text{ }\mu\text{m}$ O_2 nightglow emission line lies very near to the $1.31\text{ }\mu\text{m}$ spectral transparency window visible in night side thermal emission. Bands present in both tables correspond to those that can be used to sense both the day and night of Venus. The types of possible spectroscopic and imaging observations are also specified, as well as the physical parameters that can be inferred (dynamics, composition, thermal structure...) and a constraint for their inferred vertical distribution. Introductory references are also provided, with priority for detailed reviews, ground-based studies, and/or most recent measurements (including spacecraft). Fig. 1 displays most of the information from Tables 1 and 2 in the framework of the spectral coverage Akatsuki's cameras (Nakamura et al., 2007). Even though the atmospheric transmission on the Earth severely limits the observations at certain spectral ranges, these limitations can be partially overcome by airborne observations, balloons and space telescopes, and also using sophisticated radiative transfer models able to remove from the Venus spectra blended with lines originated by the terrestrial atmosphere (Cotton et al., 2012).

Table 1 displays the relevant bands when observing the day side of Venus, doubtless the easiest target for amateur observers since ultraviolet (UV), visible (VIS) and near-infrared (NIR) commercial filters allow to perform wind measurements with feature tracking at the upper clouds (Mousis et al. 2014). Some IR CO_2 bands can be also used on the day side to infer the altitude of the cloud tops (Ignatiev et al., 2009), and valuable observations can be done with the new generation of high-resolution spectrometers, from measuring Doppler winds (Machado et al., 2012) to studying the sulphur cycle (Marcq et al., 2011), the unknown UV ab-

sorber whose absorption might extend beyond the band in Table 1 (Tomasko et al., 1980; Marcq et al., 2011), or the thermal structure and complex atmospheric circulation of the upper atmosphere with the non-LTE emissions (Sornig et al., 2008; Peralta et al., 2016). Unfortunately, observations at FUV—which allow to study the ion population and circulation in the Venus' thermosphere—and at some infrared ranges are impossible from the ground because of Earth's low atmospheric transmittance; alternatively, airborne or satellite observations should be used (Bougher et al., 2006).

Table 2 exhibits the wavelengths useful to sense the night side of Venus. These are dominated by bands in the infrared range usually demanding longer exposures and more sensitive cameras, although amateur observers have observed the thermal emission from the Venus surface and atmosphere using NIR filters (Mousis et al. 2014). In addition to several types of nightglow occurring above the cloud tops at UV and VIS wavelengths (Gray et al., 2014), some critical CO_2 bands can be utilized to derive the atmospheric temperature at multiple levels (Grassi et al., 2008) or the altitude of the cloud tops (Ignatiev et al., 2009). Of great relevance is the IR interval $2.30\text{--}2.47\text{ }\mu\text{m}$, which allow to derive multiple compounds from the deep atmosphere below the clouds (Arney et al., 2014). IR wavelengths at $1.74\text{ }\mu\text{m}$ and $2.32\text{ }\mu\text{m}$ can be also used to sense the motions of the lower clouds (Limaye et al., 2006; Sánchez-Lavega et al., 2008). Since the clouds of Venus can potentially generate lightning, predicted bands to observe them are also included (Dubrovin et al., 2010; Pérez-Invernón et al., 2016). Although out of the range of our study, we would like to mention that X-ray observations have also been reported on the nightside of Venus (Dennerl et al., 2002; Afshari et al., 2016).

Table 1

Venus's spectral bands for day side observations. Wavelength ranging from FUV to microwave are expressed in nanometres (nm), micrometres (μm) and Gigahertz (GHz) depending on the spectral region, and these are followed by an asterisk (e.g. μm^*) when the band can be used for the same purpose in both day and night sides of Venus. The technique can be imaging (IMAG) or spectroscopy (SPEC), while the products are the parameters that can be derived are gas abundances (e.g. $[\text{CO}_2]$), atmospheric temperature (Temp), aerosols' properties and distribution (Aer), winds and waves from feature tracking (Track), winds from the Doppler technique (Doppler), and surface studies (mapping, temperature, composition and geological processes) are expressed as Surf. The vertical range sensed, processes involved in the emerging radiance and References are also specified.

| Wavelength | Products | Altitude | Processes | References |
|---|--|---------------------------------------|--|---|
| 80–140 nm (esp. 83.4, 130.4, 135.6, 130.4, 135.6) | SPEC: [OI], [OII], [O/CO ₂], Temp, Doppler? | 130–250 km | O airglow | Bougher et al. (2006), Hubert et al. (2010), Masunaga et al. (2015) |
| 121 nm* | SPEC: [H], [D/H] | 100–8000 km | Hydrogen Lyman- α emission | Chaufray et al. (2015), Bertaix (1989) |
| 170–320 nm | SPEC: [SO ₂], Aer IMAG: Track | > 70 km | SO ₂ absorption + sunlight reflected by clouds | Jessup et al. (2015) |
| 350–460 nm (max. contrast at 360–370 nm) | SPEC: [UV absorber], Doppler, Aer IMAG: Track, Aer | 63–71 km | UV absorption + sunlight reflected by clouds | Tomasko et al. (1980), Sánchez-Lavega et al. (2008), Mousis (2014) |
| 570–680 & 900–1000 nm | SPEC: Aer, Doppler IMAG: Track, Aer | 55–65 km | VIS & NIR sunlight reflected by clouds | Sánchez-Lavega et al. (2008), Mousis (2014), Hueso et al. (2015), Takagi and Iwagami (2011) |
| 1.27 μm | SPEC: [O] IMAG: [O], Track | ~ 95 km | O ₂ airglow | Connes et al. (1979) |
| 1.20, 1.44, 1.51–1.70, 2.02 μm | SPEC/IMAG: Altitude of clouds' top, Track | 60–75 km | CO ₂ absorption | Ignatiev et al. (2009), Takagi and Iwagami (2011) |
| 2.59–2.65, 3.44, 4.53, 4.75 μm | SPEC: [H ₂ O], [HDO], [HF], [CO ₂], [HCl], [CO], Temp | 70–74 km | Absorption by several species & isotopologues | Bjoraker et al. (1992), Krasnopolsky (2010, 2014), Marcq et al. (2015) |
| 4.20–4.40 μm 10.423 μm | SPEC: Temp, Doppler IMAG: Track | 105–135 km | CO ₂ nLTE emission | Peralta et al. (2016), Sornig et al. (2008) |
| 4.75 μm | SPEC: [CO], Temp. | 104–111 km | CO nLTE emission | Krasnopolsky (2014) |
| 5.3 μm^* | SPEC: [NO] | From 100 km to below 50 km | NO rotational-band emission | Krasnopolsky (2006) |
| 7.35 μm^* | SPEC: [HDO], [SO ₂], Temp | 60–70 km | Thermal emission + sunlight reflected by cloud tops + emission HDO + SO ₂ | Encrenaz et al. (2012) |
| 5.00–... μm^* 8.40–12.96 μm^* | SPEC: Temp IMAG: Temp, Track | 60–70 km | Thermal emission + sunlight reflected by cloud tops | García-Muñoz et al. (2013), Sato (2014), Takagi and Iwagami (2011) |
| 13.0–13.2 μm^* 13.6–13.9 μm^* | SPEC: Temp IMAG: Temp, Track | 69–75 km 75–81 km | CO ₂ thermal emission | Taylor et al. (1980) |
| 14.6–14.8 μm^* | | 80–96 km | | |
| 893, 335, 226 GHz* | SPEC: [HDO] | 65–100 km | HDO absorption | Encrenaz et al. (2015), Sandor (2005), Hartogh et al. (2014) |
| 625 GHz* | SPEC: [HCl] | 70–100 km | HCl absorption | Sandor (2012), Hartogh et al. (2014) |
| 115, 230, 345 GHz* | SPEC: [CO], Doppler, Temp | 90–110 km (different for day & night) | CO absorption | Lellouch et al. (2008), Clancy et al. (2012) |
| 352.8 GHz* | SPEC: [ClO] | >85 km | ClO absorption | Sandor (2013) |
| 346.6 GHz* | SPEC: [SO ₂], [SO ₃] | ~80–100 km | SO _x absorption | Encrenaz et al. (2015) |
| 23–1.385 GHz* | SPEC: [SO ₂], [H ₂ SO ₄], Temp | 50–60 km | Thermal emission CO ₂ CIA (main) with opacity by SO ₂ & H ₂ SO ₄ | Butler et al. (2001), Jenkins et al. (2002) |
| 15–1.50 GHz* | IMAG: Surf | Surface | Surface Radio-Thermal emission | Campbell et al. (2015) |

Finally, both Tables exhibit at longer wavelengths several common bands of special interest, since they allow to visualize the day and night of Venus simultaneously and at similar heights, except for concrete exceptions like the CO (Lellouch et al., 2008). Mid-infrared bands sense the thermal emission from the upper clouds to about 90 km (Taylor et al. 1980; García-Muñoz et al., 2013), while sub-millimetre and microwave heterodyne observations can cover a wider vertical range in the upper atmosphere and allow to sense the rotational transitions of mesospheric minor species such as CO, HDO, SO, SO₂, and HCl (Encrenaz et al., 2015). CO lines are used to sense the atmospheric temperature and Doppler winds using sub-millimetre heterodyne spectroscopy (Lellouch et al., 2008; Clancy et al., 2012). At longer wavelengths in microwave, we can sense the atmosphere below the clouds since the CO₂ collision-induced absorption becomes weaker (Jenkins et al., 2002). Lightning was indirectly detected via whistler waves at 1–10 Hz by Venus Express (Russell et al., 2007), and it has been also proposed that it might be detected at 5–50 MHz using UHF

detectors in a low Venus orbit (Majid et al., 2013). At further long wavelengths in radar domain (~2.3 GHz, 12.6 cm), the atmosphere of Venus becomes quite transparent and the surface can be imaged (Campbell et al., 2015), while subsurface sounding has been also suggested to be possible at the 10–30 MHz range (Biccardi et al., 2004).

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Table 2

Venus's spectral bands for night side observations. Information and notation is the same as in Table 1.

| Wavelength | Products | Altitude | Processes | References |
|---|--|--|---|--|
| 121 nm* | SPEC: [H], [D/H] | 100–10,000 km | Hydrogen Lyman- α emission | Chaufray et al. (2015), Bertaux (1989) |
| 190–300 nm 120–280, 208, 250–450 nm | SPEC: [N], [O] IMAG: lightning? | 102–125 km Cloud region or above? | NO nightglow Electric discharge in N ₂ –CO ₂ mixture | Krasnopolksy (2006) Dubrovin et al. (2010), Pérez-Invernón et al. (2016) |
| 400–650 nm 557.7 & 777.3 nm | SPEC: [O ₂] IMAG: Track SPEC: [O] IMAG: lightning? | 100–130 km Upper atmosphere | Herzberg II O ₂ nightglow O nightglow & aurora & electric discharge in N ₂ –CO ₂ mixture | García-Muñoz et al. (2009b) Gray et al. (2014), Dubrovin et al. (2010), Pérez-Invernón et al. (2016) |
| 850 & 900 nm 1.01 μ m | IMAG: Surf | Surface (main source) | Surf. thermal emission + clouds' opacity | Baines (2000), Mousis (2014) |
| 1.14–1.19, 1.10 μ m | IMAG: Surf, [H ₂ O], [HDO] | 0–25 km | Surf. thermal emission + Abs. HDO+H ₂ O | Bailey et al. (2008b), Bézard et al. (2011) |
| 1.224 μ m 1.269 μ m | SPEC: [NO] SPEC/IMAG: [O], Temp, Track | 95–132 km 95–105 km | NO nightglow O ₂ airglow from O recombination | García-Muñoz et al. (2009a) Ohtsuki et al. (2008), Bailey et al. (2008a), Connes et al. (1967), Krasnopolksy (2010) |
| 1.273–1.304 μ m | SPEC: [HF] | Cloud tops | HF absorption | Mueller et al. (2008) |
| 1.31 μ m | SPEC: ? IMAG: Track? | Between clouds and surface | Scattering of surface thermal emission, weak | |
| 1.40–1.49 μ m 1.51 & 1.55 μ m | SPEC: [OH] SPEC: [H ₂ O] IMAG: [H ₂ O] | ~ 96 km 20–35 km | OH nightglow Unexplained minimum in H ₂ O absorption | Piccioni et al. (2008) Wilson et al. (2009) |
| 1.58 μ m 1.727–1.758 μ m | SPEC/IMAG: [O], Temp? SPEC: [HCl], [H ₂ O] IMAG: Track | Upper atmosphere Species: 13–34 km Clouds: 44–48 km | O ₂ airglow Absorption by several species + Thermal Emission + Clouds opacity | Bailey et al. (2008b) Sánchez-Lavega et al. (2008), Bézard (2007), Arney et al. (2014) |
| 2.30–2.47 μ m | SPEC: [CO], [HDO], [H ₂ O], [OCS], [SO ₂] IMAG: Track | Species: 26–40 km Clouds: 44–48 km | Absorption by several species + Thermal Emission + Clouds opacity | Limaye et al. (2006), Bézard (2007), Arney et al. (2014), Takagi and Iwagami (2011) |
| 2.60–3.14 (2.80 & 2.94) 4.25–5.00 4.53 μ m 5.3 μ m* 7.35 μ m* | SPEC: [OH] SPEC: Temp SPEC: [CO] SPEC: [NO] SPEC: [HDO], [SO ₂], Temp | 95 ± 10 km 50–100 km ~ 70 km From 100 km to below 50 km 60–70 km | OH nightglow CO ₂ emission CO absorption NO rotational-band emission Thermal emission from upper clouds + emission HDO + SO ₂ | Piccioni et al. (2008) Grassi et al. (2008) Marcq et al. (2015) Krasnopolksy (2006) Encrenaz et al. (2012) |
| 8.25 μ m | IMAG: lightning? | Cloud region or above? | Electric discharge in N ₂ –CO ₂ mixture | Pérez-Invernón et al. (2016) |
| 3.68–3.94 μ m 4.81–4.89 μ m 5.00–... μ m* 8.40–13.0 μ m* | SPEC: Temp IMAG: Temp, Track | 60–70 km | Thermal emission from upper clouds | Taylor et al. (1980), Sato (2014), García-Muñoz et al. (2013), Takagi and Iwagami (2011) |
| 13.0–13.2 μ m* 13.6–13.9 μ m* 14.6–14.8 μ m* 136.1 μ m | SPEC: Temp IMAG: Temp, Track IMAG: lightning? | 69–75 km 75–81 km 80–96 km Cloud region or above? | CO ₂ thermal emission Electric discharge in N ₂ –CO ₂ mixture | Taylor et al. (1980) Pérez-Invernón et al. (2016) |
| 893, 335, 226 GHz* | SPEC: [HDO] | 65–100 km | HDO absorption | Encrenaz et al. (2015), Sandor (2005), Hartogh et al. (2014) |
| 625 GHz* | SPEC: [HCl] | 70–100 km | HCl absorption | Sandor (2012), Hartogh et al. (2014) |
| 115, 230, 345 GHz* | SPEC: [CO], Doppler, Temp | 90–110 km (different for day & night) | CO absorption | Lellouch et al. (2008), Clancy et al. (2012) |
| 352.8 GHz* 346.6 GHz* 23–1.385 GHz* | SPEC: [ClO] SPEC: [SO], [SO ₂] SPEC: [SO ₂], [H ₂ SO ₄], Temp | > 85 km ~80–100 km 50–60 km | ClO absorption SO ₂ absorption Thermal emission CO ₂ CIA (main) with opacity by SO ₂ & H ₂ SO ₄ | Sandor (2013) Encrenaz et al. (2015) Butler et al. (2001), Jenkins et al. (2002) |
| 15–1.50 GHz* | IMAG: Surf | Surface | Surface Radio-Thermal emission | Campbell et al. (2015) |

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